

On the importance of setting 3-D stress field in simulations of on- and off-fault deformation

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SUMMARY

During the last decades, many numerical models have been developed to investigate the conditions for seismic and aseismic slip. Those models explore the behaviour of frictional faults, embedded in either elastic or inelastic media and submitted to a far field loading (seismic cycle models), or initial stresses (single dynamic rupture models). Those initial conditions impact both on-fault and off-fault dynamics. Because of the sparsity of direct measurements of fault stresses, modellers have to make assumptions about these initial conditions. To this day, Anderson's theory is the only framework that can be used to link fault generation and reactivation to the 3-D stress field. In this work, we look at the role of the 3-D stress field in modelling a 2-D strike-slip fault under plane-strain conditions. We show that setting up an incorrect initial stress field, based on Anderson's theory, can lead to underestimation of the damage zone width by up to a factor of six, for the studied cases. Moreover, because of the interactions between fault slip and off-fault deformation, initial stress field influences the rupture propagation. Our study emphasizes the need to set up the correct initial 3-D stress field, even in 2-D numerical simulations.

Key words: Elasticity and anelasticity; Friction; Numerical modelling; Earthquake dynamics; Dynamics and mechanics of faulting.

1 INTRODUCTION: MODELLING FAULT SLIP

Catastrophic, seismic events of large magnitude ($M_w > 7$) remain relatively rare, with a recurrence time of several decades up to a millennium (Cubas *et al.* 2015). As a consequence, observations are sparse and numerical models are powerful tools to explore the conditions for seismic and aseismic fault slip. Numerical models that account for both seismic slip and long-term slow slip are challenging because of the wide range of temporal and spatial scales involved (Lapusta *et al.* 2000). Hence, compromises are made to reduce the computational cost depending on the goal of the model. These compromises can be safely categorized as models that focus on a single dynamic rupture event and ones that model multiple seismic cycles.

Single dynamic rupture models were some of the first analytical and numerical models developed for earthquakes. They reproduce an event from the moment it turns dynamic to its arrest and have provided important insights into earthquake mechanics (Kostrov 1964; Andrews 1976; Madariaga 1976; Das & Kostrov 1986, 1987, among many others that followed).

For seismic cycle models, the focus is on integrating all stages of fault slip (inter-, co- and post-seismic) over thousands of years (Erickson *et al.* 2020). This is indeed critical as pre-stress inherited

from aseismic slip and prior seismic events likely play a determinant role on where earthquakes will nucleate and how far their rupture will propagate (Thomas *et al.* 2014). The most popular strategy is to model a single planar fault governed by rate-and-state (R&S) friction law, embedded in a purely elastic medium. (e.g. Ben-Zion & Rice 1997; Lapusta *et al.* 2000; Richards-Dinger & Shearer 2000; Kato 2004; Barbot *et al.* 2012; Im *et al.* 2020; Liu *et al.* 2020). In these models, frictional heterogeneities are invariably advocated to reproduce the full slip spectrum (creep, slow slip events, earthquakes, etc.). More recent models also account for complex fault geometry (Romanet *et al.* 2018; Ozawa & Ando 2021; Uphoff *et al.* 2022, among others). However, due to timescales that vary over several decades of orders of magnitude, the most popular compromise that is made is to ignore the wave mediated stress transfer (inertial dynamics) and only account for linear elastic static stress transfer along with the instantaneous, local traction contribution. Such class of models are called quasi-dynamic models.

Natural fault zones, however, that is the fault plane and its surrounding medium, are intricate structures in constant evolution in response to tectonic loading. As an example, during an earthquake, part of the stored elastic strain energy is dissipated in off-fault deformation, or damage, which in turn radiate and affect the slip dynamics of the main fault. If we provide an idealized description, fault zones comprise of a non-planar fault core, where most of the

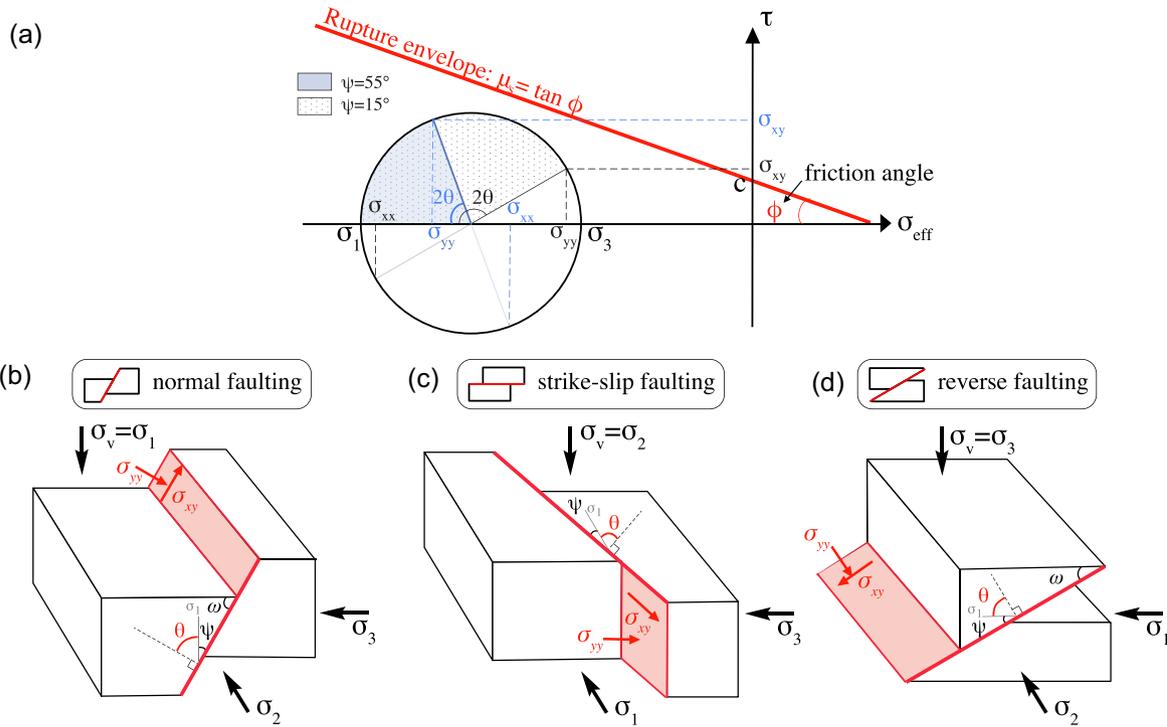


Figure 1. (a) Mohr–Coulomb criterion in (σ, τ) -space. (b), (c) and (d) Orientation of the failure plane with respect to the largest principal stress for normal, strike-slip and reverse faulting, respectively. The angle ϕ correspond to the slope of the failure envelope on a Mohr diagram and defines the static friction $\mu = \tan \phi$. The remaining angles Ψ , θ and ω are defined within the plane perpendicular to σ_2 . Ψ and θ corresponds to the angle between σ_1 and the fault plane (Ψ) or, following the Mohr–Coulomb convention, the normal to the fault plane (θ). ω gives to the dip angle of the fault.

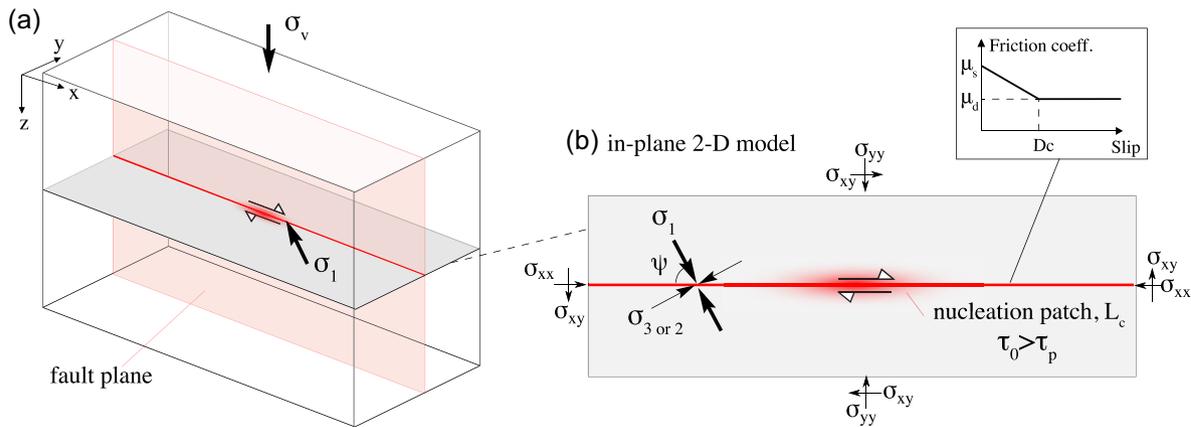


Figure 2. Modelling set up for a dynamic rupture on a strike-slip fault. (a) Schematic of the fault zone in 3-D. (b) Zoom on the 2-D plane that hosts the modelled rupture. The orientation of maximum compressive stress σ_1 is set at an angle Ψ to the fault plane that is governed by slip weakening friction law. The nucleation patch is set by increasing the initial shear stress σ_{xy} slightly above the fault strength $\mu_s(-\sigma_{yy})$ over a length L_c .

displacement has occurred, surrounded by a damage zone that has a spatial scale of the order of meters to kilometres (e.g. Chester *et al.* 1993; Biegel & Sammis 2004; Faulkner *et al.* 2011). Frequently, the fault core corresponds to an extremely narrow band and the damaged wall rock includes layers of gouge and breccia bordered by fractured rocks. The last two layers are included in the damage zone because they lacked extensive shearing. These structures are of key importance in the mechanics of faulting. For example, fault roughness has an effect on the fault resistance to slip, that is the fault strength (e.g. Dunham *et al.* 2011a; Tal & Faulkner 2022). Laboratory experiments of earthquakes in a damaged medium show that there is an intimate interaction between the rupture and off-fault

damage zone (Sammis *et al.* 2009; Bhat *et al.* 2010; Biegel *et al.* 2010). The density of this damage has a direct impact on the elastic moduli of the bulk (Walsh 1965b, a; Faulkner *et al.* 2006), therefore, on the quantity of strain energy which is stored and further released by fault slip. In fact, systematic micro- and macrostructural field studies have been performed on damage zones (e.g. Shipton & Cowie 2001; Manighetti *et al.* 2001, 2004; Dor *et al.* 2006; Mitchell & Faulkner 2009; Savage & Brodsky 2011; Johnson *et al.* 2021; Rodriguez Padilla *et al.* 2022), as a key component to understand the energy balance of earthquakes (e.g. Rice 2002; Kanamori 2006; Okubo *et al.* 2019). Off fault damage is observed from the mesoscopic scale to the microscopic scale, with a microfracture density

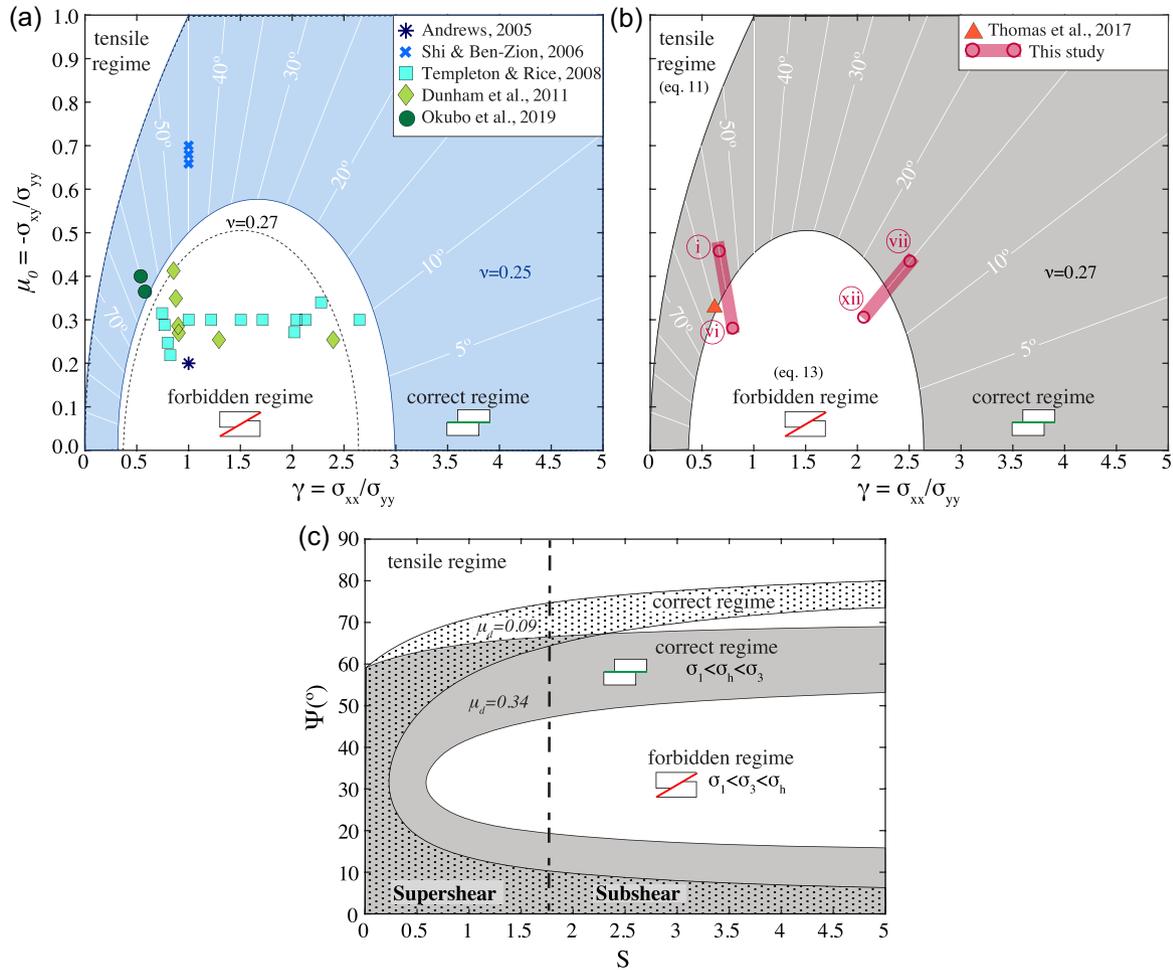


Figure 3. Criterion for accurate initial stress field under plane-strain approximation displayed as $\gamma - \mu_0$, plots (a) and (b) and $\Psi - S$ for plot (c). The shaded areas represent the conditions for which the initial stresses favour a strike-slip motion, that is when eq. (14) is satisfied. In the $\gamma - \mu_0$ plot, this criterion depends on ν (0.25 and 0.27 for plot (a) and (b), respectively), but is independent from other parameters. Contours lines shows Ψ values. Colour dots shows the initial parameters for published simulations of strike-slip faulting with plane-strain approximation. In plot (b), the shaded red areas shows the two sets of parameters explored in this study, with $\Psi = 15^\circ$ and $\Psi = 55^\circ$. The red empty circles correspond to the end members. In subplot (c) the criterion is represented as a function of S and Ψ for $\mu_s = 0.6$. Plotting the area corresponding to an accurate initial stress depends on a third parameter, here given by μ_d . We plot the criterion for two cases, $\mu_d = 0.09$ (dotted area) and $\mu_d = 0.34$ (grey area).

Table 1. Constants used in all models.

Parameter	Symbol	Value
Depth	z	-2.5 km
Material density	ρ	2700 kg m ⁻³
S -wave velocity	c_s	3120 km s ⁻¹
P -wave velocity	c_p	5600 km s ⁻¹
Poisson's ratio	ν	0.27
Characteristic slip	D_c	1 m
Pore pressure coefficient	λ	0.4

that decreases exponentially away from the fault core (Mitchell & Faulkner 2009). The width of the damage zone is also believed to be decreasing with depth, forming a ‘flower-like structure’ (Ben-Zion et al. 2003; Cochran et al. 2009). However Okubo et al. (2019) have demonstrated numerically that, despite its reduction in spatial-extent with depth, energetically speaking, the contribution of the off-fault damage increases with depth.

Thus, as oppose to seismic cycle codes, a second set of models have been developed to catch the full slip dynamics, the wave propagation and the interactions with the off-fault medium during an earthquake. With these models, researchers have explored the effect of complex fault geometry and/or the effect of off-fault damage on seismic rupture. Some models treat the bulk as a solid linear-elastic material and prescribe a low-velocity zone around the fault to account for damage (e.g. Cappa et al. 2014; Huang et al. 2014). Another set of models has explored the effect of spontaneous dynamic generation of off-fault deformation using a Mohr–Coulomb (e.g. Andrews 2005; Ben-Zion & Shi 2005; Hok et al. 2010; Gabriel et al. 2013) or Drucker–Prager (e.g. Templeton & Rice 2008; Ma 2008; Dunham et al. 2011a; Johri et al. 2014) based plastic constitutive laws. Another class of models have treated off-fault damage as tensile cracks, using a stress- (Yamashita 2000) or fracture-energy-based (Dalguer et al. 2003) criterion. Okubo et al. (2019) used a finite discrete element method (FDEM) that allows spontaneous nucleation and propagation of off-fault fracture network in a medium. These studies have provided a good insight on the effect of a fault zone structure on dynamic ruptures but the models do not account

Table 2. Initial parameters used for the two end-members of each set of models.

	Parameter	Symbol	Initial stress field			
			Case (i) Strike-slip	Case (vi) Reverse	Case (vii) Strike-slip	Case (xii) Reverse
Input	Angle	Ψ	55°	55°	15°	15°
	Seismic ratio	S	2	2	1	1
	Stress drop (MPa)	$\Delta\tau$	10.2	10.2	10.6	10.6
	Static friction	μ_s	0.68	0.52	0.68	0.52
Resulting parameters	Dynamic friction	μ_d	0.34	0.16	0.19	0.09
	Full stress tensor (MPa)	σ_{yy}	-91	-84	-43	-50
		σ_{xx}	-61	-67	-108	-102
		σ_{zz}	-42	-42	-42	-42
		σ_{xy}	41	23	19	15
		Principal stresses (MPa)	σ_1	-120	-101	-113
		σ_2	-42	-51	-42	-45
		σ_3	-32	-42	-38	-42
	Process zone (m)	R_0	1047	1047	1570	1570
Nucleation length (m)	L_c	6562	6562	4374	4374	

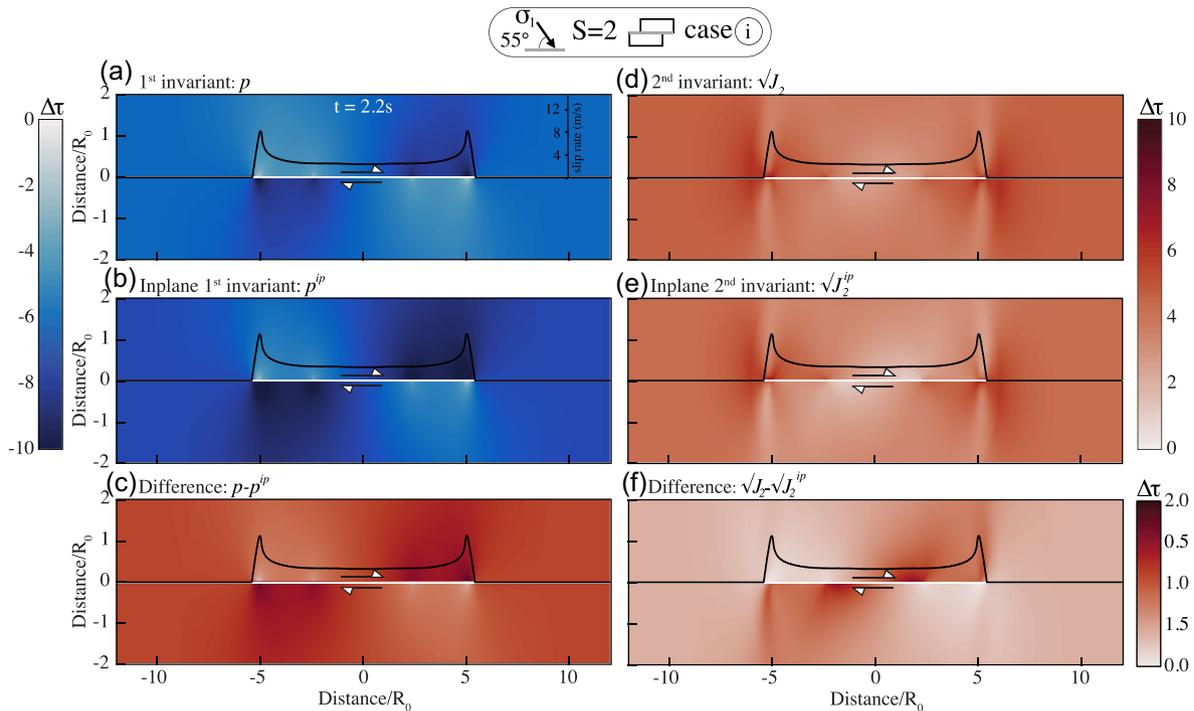


Figure 4. First invariant of the stress tensor (a, b) and second invariant of the deviatoric stress tensor (d, e) computed with the 3-D stress tensor (a, d), the in-plane stress tensor (b, e), at $t = 2.9$ s for case (i) with $S = 2$ and $\Psi = 55^\circ$. Panels (c, f) give, respectively, the difference between the way of computing the first invariant of the stress tensor and the second invariant of the deviatoric stress tensor. Invariants are normalized by the dynamic stress drop $\Delta\tau$ (eq. 19). Slip rate on the fault (black curves) is superimposed on the graphs.

for the observed coseismic change of elastic properties in the bulk (Hiramatsu *et al.* 2005; Brenguier *et al.* 2008; Froment *et al.* 2014) which also impacts the rupture dynamics. This can be achieved by using a homogenized damage mechanics based constitutive law (e.g. Lyakhovsky *et al.* 1997; Bhat *et al.* 2012; Xu *et al.* 2014; Thomas *et al.* 2017b; Thomas & Bhat 2018). A vast majority of the investigations cited above are done under 2-D plane-strain conditions.

Finally, some models have been developed to look at the zeroth order effect of the fault core (as opposed to a fault interface) or the damage zone on the seismic cycle. To overcome the computational cost, they have to compromise on both the fault slip dynamics and the dynamics of bulk evolution (off-fault crack growth). As an example, van den Ende *et al.* (2018) have represented the

fault core by a shear band but the bulk remains elastic. Kaneko *et al.* (2011); Idini & Ampuero (2020), and Abdelmeguid *et al.* (2019) among others, looked at the effect of a prescribed low velocity fault zones (LVFZ), but, by construction, the bulk is still a passive elastic body. They used quasi-static or quasi-dynamic approximations to solve the problem. In their model, Thakur & Huang (2021) imposed (thus not driven by the model) a time-dependent shear modulus evolution to account for coseismic drop and post-seismic recovery of elastic moduli. Erickson *et al.* (2017) has applied quasi-dynamic analysis to explore the effect of plasticity throughout the earthquake cycle. Preuss *et al.* (2020) have developed a 2.5-D model with a viscoelasto-plastic crust subjected to rate- and state-dependent friction to model conjointly the rapid

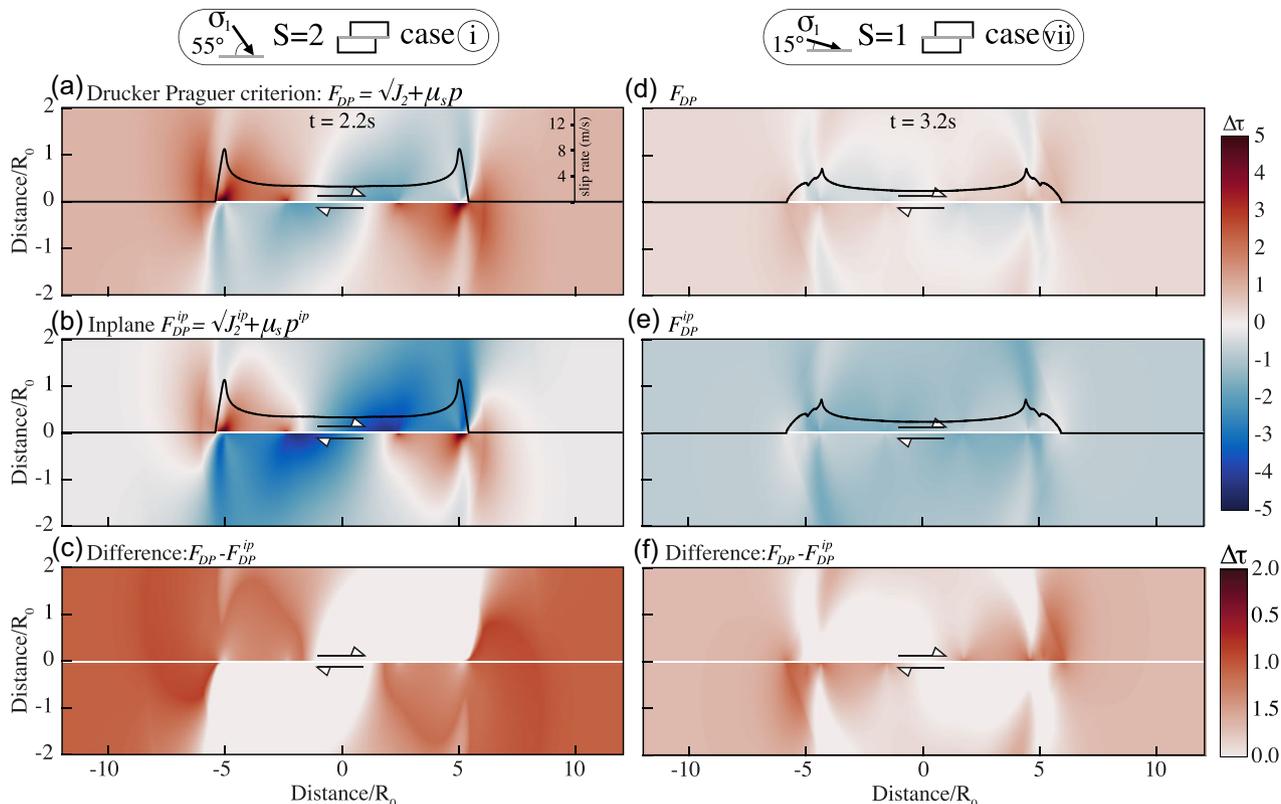


Figure 5. Drucker–Prager criterion for case (i) with $S = 2$ and $\Psi = 55^\circ$ (a, b) and for case (vii) with $S = 1$ and $\Psi = 15^\circ$ (d, e), at $t = 2.2$ s. We compute the invariants accounting for σ_{zz} (a, d), or only using the in-plane stress field (b, e). We show the differences $F_{DP} - F_{DP}^{ip}$ in panels (c, f) for cases (i) and (vii), respectively. The difference is computed for the positive values only. Different versions of the Drucker–Prager criteria and their differences are normalized by the dynamic stress drop $\Delta\tau$ (eq. 19). Slip rate on the fault (black curves) is superimposed on the graphs.

deformation in the elastic–brittle upper crust and the relaxation in the deeper viscoelastic crustal substrate and their influence on each other.

In all cases, regardless of the end goal, setting the initial and boundary conditions (initial stresses for a single rupture, far field loading for seismic cycles) directly impacts both on-fault and off-fault dynamics. But what constraints do modellers have on the stress state of a fault? This will be discussed in the section below, followed-up by a review on how boundary conditions are set up in the community. We will then present a simple method to accurately define the initial stress field for in-plane conditions. To demonstrate its importance, we will explore its influence on-fault and off-fault deformation. Results are summarized and discussed in the last section.

2 STRESS FIELD AND FAULTING

2.1 The Anderson theory

In the upper crust, large strains are accommodated by fault systems either seismically or aseismically. It either leads to the formation of new fractures in the crust, or reactivates previously existing faults. Fault systems evolve and acquire their general geometry by the progressive amalgamation of such fractures (Cowie & Scholz 1992). *In situ* measures of stresses on a fault, at any location, at any time are impossible to achieve, hence it has to be approached theoretically. A connection between the geometry of fault systems and the forces that formed them was first proposed by (Anderson 1905). His theory

related the initial formation of faults to the state of stress in the crust, under the assumption that rocks are isotropic, homogeneous and intact. This theory is based on the mathematical result that at every point of a stressed rock, three planes can be found on which no shear traction acts. Those planes are perpendicular to one another and are called the principal planes. The corresponding stresses acting along the three principal directions are called the principal stresses. By convention, the stresses are assumed to be positive in tension. Therefore, in a compressive regime, we have the following inequality :

$$\sigma_1 < \sigma_2 < \sigma_3. \quad (1)$$

Then Anderson assumed that, apart from those three planes, a plane with maximum tangential stress exists, on which the faulting should initiate. For brittle shear failure, and for a fluid saturated rock mass with pore fluid pressure P_f (with $P_f > 0$), this Coulomb failure criterion may be written as:

$$\tau_y = c + \mu(-\sigma_{\text{eff}}) = c + \mu(-\sigma_n - P_f), \quad (2)$$

where σ_n is the normal stresses acting on the plane and $\sigma_{\text{eff}} = \sigma_n + P_f$ corresponds to the effective normal stress. The variable c is the cohesion and $\mu = \tan\phi$ is the coefficient of friction. The angle ϕ corresponds to the slope of the failure envelope on a Mohr diagram (Fig. 1) and θ , on the same diagram, corresponds to the angle between the normal to the fault plane and σ_1 (Fig. 1). Hence

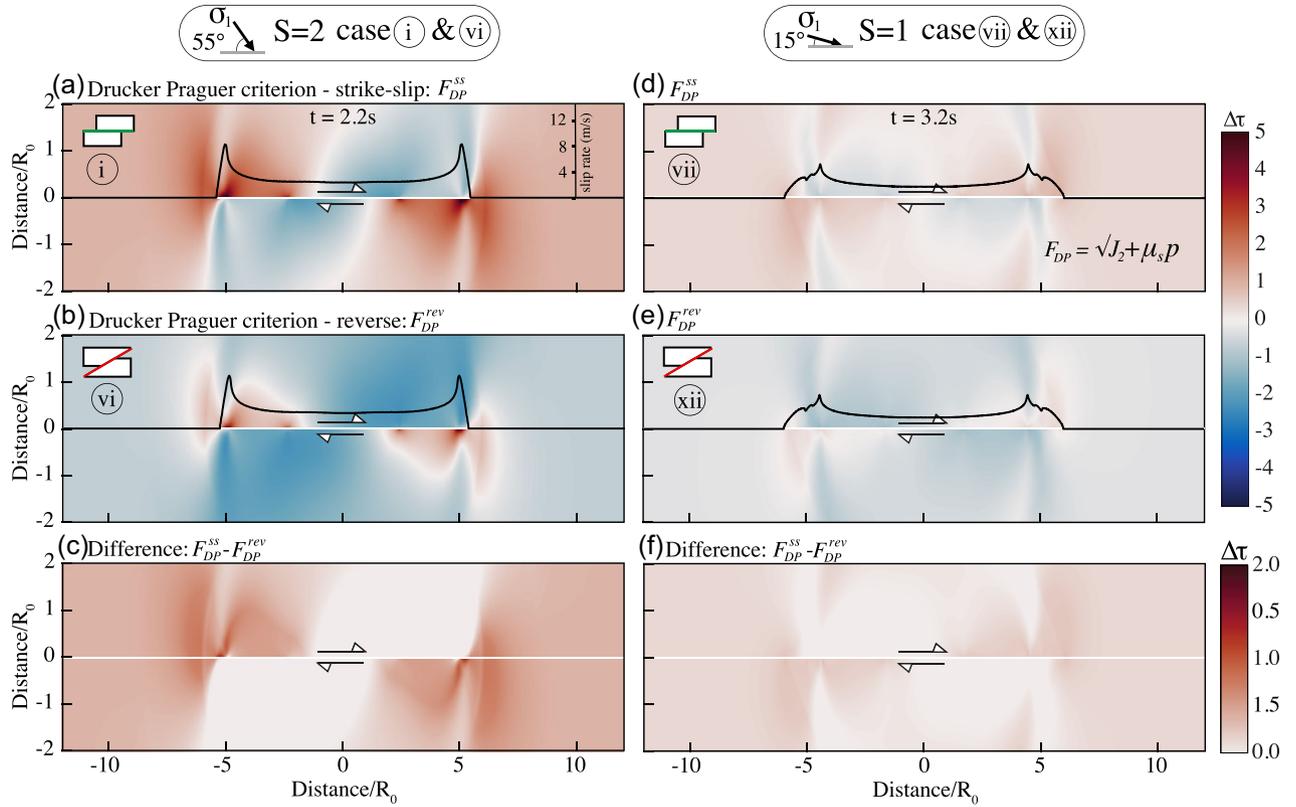


Figure 6. Drucker–Prager criterion for cases (i) and (vi), with $S = 2$ and $\Psi = 55^\circ$ (a, b) and for cases (vii) and (xii), with $S = 1$ and $\Psi = 15^\circ$ (d, e). Cases (i) and (vii) correspond to an initial strike-slip stress field (a, d) and cases (vi) and (xii) to an initial reverse stress field (b, e). Panels (c, f) give the difference between cases (i) and (vi) and between cases (vii) and (xii), respectively. The four models are plotted against the criterion for accurate initial stress in Fig. 3. Drucker–Prager criterions and differences are normalized by the dynamic stress drop $\Delta\tau$ (eq. 19). Slip rate on the fault (black curves) is superimposed on the graphs.

we can write:

$$\begin{aligned}\sigma_{yy} &= \left(\frac{\sigma_1 + \sigma_3}{2}\right) - \left(\frac{\sigma_3 - \sigma_1}{2}\right) \cos 2\theta \\ \sigma_{xx} &= \left(\frac{\sigma_1 + \sigma_3}{2}\right) + \left(\frac{\sigma_3 - \sigma_1}{2}\right) \cos 2\theta \\ \sigma_{xy} &= \left(\frac{\sigma_3 - \sigma_1}{2}\right) \sin 2\theta.\end{aligned}\quad (3)$$

The radius R of the Mohr circle is then given by:

$$R = \left(\frac{\sigma_3 - \sigma_1}{2}\right) = \sqrt{\left(\frac{\sigma_{xx} - \sigma_{yy}}{2}\right)^2 + \sigma_{xy}^2}.\quad (4)$$

In numerical studies, it is more common to define Ψ , the angle between the fault plane and σ_1 (Figs 1 and 2), such as $\Psi = \pi/2 - \theta$. Hence, as illustrated in Fig. 1(a), the relationship between the magnitudes of σ_{yy} and σ_{xx} depends on Ψ in the following manner

$$\begin{aligned}|\sigma_{yy}| &\leq |\sigma_{xx}| \text{ if } 0^\circ \leq \Psi \leq 45^\circ \\ |\sigma_{yy}| &> |\sigma_{xx}| \text{ if } 45^\circ < \Psi \leq 90^\circ.\end{aligned}\quad (5)$$

Then, for an optimal angle $\theta_{\text{opt}} = \pi/4 + \phi/2$ shear failure occurs on a plane containing the σ_2 direction, when $\sigma_{yy} = \sigma_{\text{eff}}$ and $\sigma_{xy} = \tau_y$, that is when the failure envelop is tangential to the Mohr's circle. Relying on experimental studies, the internal friction μ generally lies between 0.5 and 1 (Jaeger & Cook 1979). Hence the optimal angle θ_{opt} varies between 58° and 68° and the new fault should make an angle Ψ with σ_1 between 32° and 22° , respectively. Following

the same reasoning for a pre-existing fault plane, if we consider a static coefficient of friction $\mu_s = 0.6$, which corresponds to a large variety of rocks and minerals (Byerlee 1978), faults are optimally oriented when they make an angle $\Psi \simeq 30^\circ$ with σ_1 .

The foundation of the Anderson's theory lies in the observation that, due to Earth's free surface, it is imperative for one principal stress to be oriented subvertically, that is to be equal to σ_v . This gives rise to three fundamental stress regimes (Fig. 1) depending on whether $\sigma_v = \sigma_1, \sigma_2$ or σ_3 : reverse faulting ($\sigma_v = \sigma_3$), strike-slip faulting ($\sigma_v = \sigma_2$) and normal faulting ($\sigma_v = \sigma_1$). Keeping $\mu_s = 0.6$ as a reference value for the static friction, it is then expected to get subvertical strike-faults and a dip angle of $\omega \approx 30^\circ$ and $\omega \approx 60^\circ$ for a thrust and a normal fault respectively (Fig. 1).

Despite the simplicity of the theory, observations are in accordance with the model for strike slip faults with low cumulative displacement (Anderson 1951; Kelly *et al.* 1998). Earthquakes have also been registered on subvertical fault striking approximately 30° to the regional σ_1 with a subvertical σ_2 : the 2000 Western Tottori earthquake in Japan (Sibson *et al.* 2012; Fukuyama *et al.* 2003; Yukutake *et al.* 2007) or the 2010 M_w 7.1 Darfield earthquake in New Zealand (Sibson *et al.* 2012). Moreover, borehole measurements and induced seismicity (Townend & Zoback 2000), palaeostresses inversion (Lisle *et al.* 2006) and earthquake focal mechanisms (C el erier 2008) suggest that Andersonian state of stress prevails over large areas within the shallow crust. Of course exceptions to the theory exist too. Well-known examples, such as the San Andreas fault (Zoback

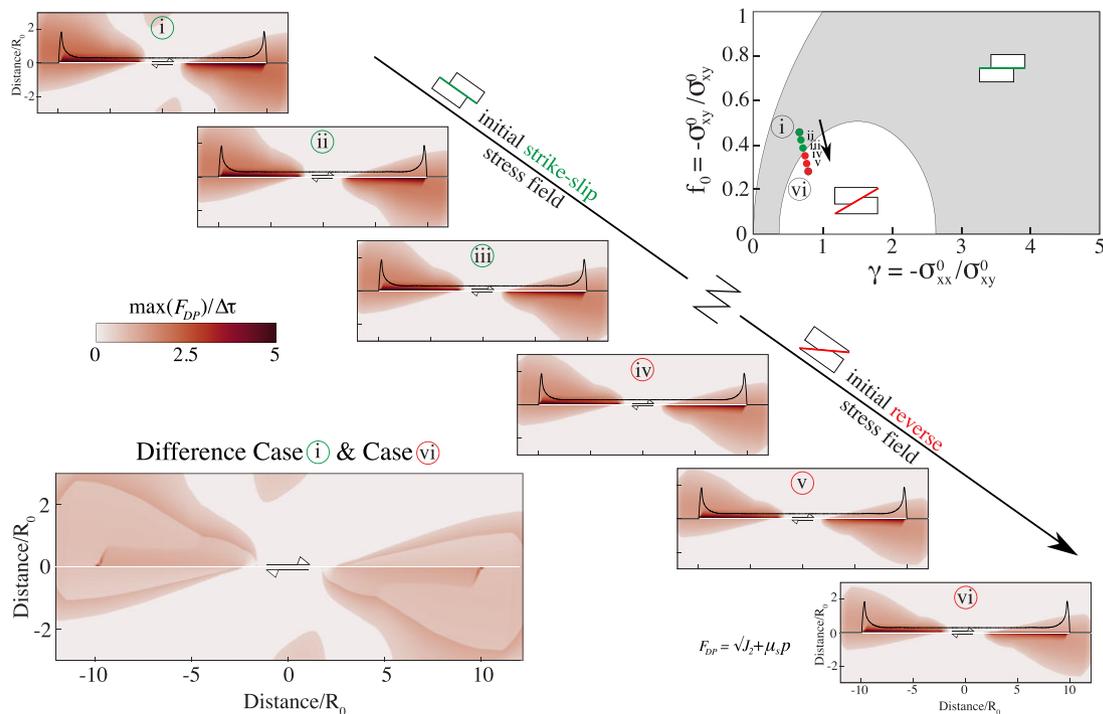


Figure 7. Drucker–Prager criterion for all models with $\Psi = 55^\circ$ and $S = 2$. We plot the maximum value F_{DP} between $t = 0$ and $t = 4$ s. Only positive values are considered (hence the colour discontinuity). We display the results from the strike-slip end-member (upper left corner) to the reverse end-member (lower right-hand corner). The initial parameters are shown in the upper right inset, in the same representation as in Fig. 3(a). The lower left inset shows the difference between the two end members. Drucker–Prager criterions and differences are normalized by the dynamic stress drop $\Delta\tau$ (eq. 19). Slip rate on the fault (black curves) is superimposed on the graphs.

et al. 1987), or the low-angle normal faults in Elba, Central Italy (Collettini & Sibson 2001) or in the Cyclades Greece (Lecomte *et al.* 2010) are mis-oriented if we consider a coefficient of friction of 0.6 (Byerlee 1978). But a lower friction on the fault plane (clay minerals) or a high pore pressure may well explain the discrepancy. If the fault essentially slips during earthquakes, weakening mechanisms such as the ones listed by Tullis & Schubert (2015) may well kick in (mineral breakdown, flash melting, thermal pressurization, etc.), leading to a very low effective coefficient of friction. Finally, polymodal faulting, with three or more sets of faults forming and slipping simultaneously are not compatible with the Anderson’s assumption that faults form parallel to the intermediate principal stress, σ_2 . Hence, excluding these exceptions, it suggests that the Anderson’s theory provides a useful framework for general considerations about fault generation and reactivation. And considering the lack of systematic, time-dependent *in situ* stress measurements, probably the only one.

2.2 Non-exhaustive review on how initial stress state is prescribed in numerical models

In single dynamic rupture models, rupture grows under the control of the prescribed initial stresses. Theoretical analyses (Poliakov *et al.* 2002; Rice *et al.* 2005; Ngo *et al.* 2012) have illustrated the effect of initial stress field on the pattern of off fault damage activation (i.e. the potential failure area, and the orientation of secondary cracks) around a propagating crack. They show that the extent and location of secondary faulting (the activated zone) is strongly affected by the orientation of principal stresses, set up by Ψ (the

angle between σ_1 and the primary fault). Steep Ψ favours inelasticity on the extensional side and shallow Ψ on the compressional side. Moreover initial stresses seem to influence the potential for rupture to follow intersecting faults with different orientation rather than the primary fault (Kame *et al.* 2003; Bhat *et al.* 2004; Fliss *et al.* 2005). Therefore, pre-stress orientation is a key parameter in numerical simulation of dynamic rupture with off-fault inelastic deformation.

Most of the numerical studies investigating the interactions between seismic rupture propagation and off fault damage usually set up a 2-D planar strike-slip fault under plane-strain conditions (Andrews 2005; Shi & Ben-Zion 2006; Templeton & Rice 2008; Dunham *et al.* 2011b; Thomas *et al.* 2017b; Okubo *et al.* 2019), such as displayed in Fig. 1. Because of the 2-D setting, the out of plane stress is often ignored when setting up the initial stress field. Usually, a normal stress and a shear stress are imposed, corresponding to two principal horizontal stresses σ_1 and σ_3 and an angle Ψ between fault and main stress direction. However, whether a fault is under a strike-slip stress field depends on the full 3-D stress field (Fig. 1). If the two in-plane stresses are σ_1 and σ_3 , then the fault is in a proper strike-slip stress field. If the two in-plane stresses are σ_1 and σ_2 , then the fault is in a reverse stress field. This would not change the slip on the fault because slip is restricted to the 2-D plane. However, the ratio between the out of plane stress and the smallest principal in-plane stress would be smaller than expected under a proper 3-D stress strike-slip stress field.

Despite the importance of initial stress field in modelling off fault deformation, the *full* 3-D initial stress field and the importance of the out-of-plane stress, in 2-D simulations, have never been discussed.

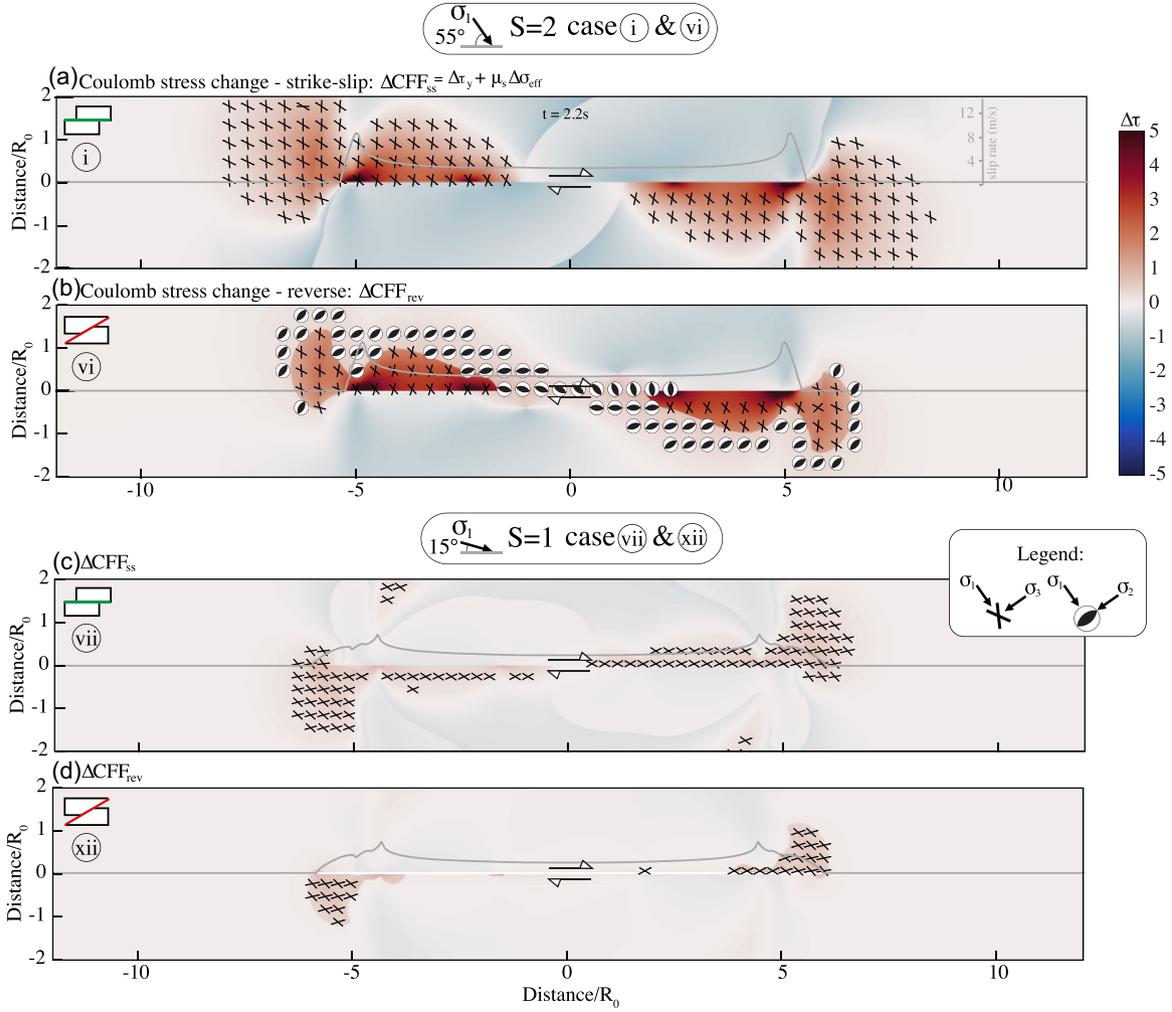


Figure 8. Coulomb stress change ΔCF due to the dynamic rupture, computed on optimally oriented planes and normalized by the dynamic stress drop $\Delta\tau$ (eq. 19), for cases (i) and (vi), with $S = 2$ and $\Psi = 55^\circ$ (a, b) and for cases (vii) and (xii), with $S = 1$ and $\Psi = 15^\circ$ (c, d). Cases (i) and (vii) correspond to an initial strike-slip stress field (a, c) and cases (vi) and (xii) to an initial reverse stress field (b, d). Superimposed on this snapshot are the conjugates planes that give a maximum value of ΔCF . The symbols inform on the expected mode of rupture: two crossed lines represent the two conjugate planes for strike-slip ruptures, the beach ball with a dark centre gives the orientation of the conjugate planes for reverse faulting, following the classic seismological convention. Those symbols are displayed in the areas where $\Delta\text{CF} > 0.2$. Slip rate on the fault (grey curve) is superimposed on the graphs.

2.3 Criterion for accurate initial stress field in 2-D plane-strain simulations

When considering a planar, strike-slip fault under plane-strain condition (Fig. 2). The initial stress state, σ_{ij} , (tensile positive) is represented as,

$$\bar{\sigma} = \sigma_{ij} \equiv \begin{pmatrix} \sigma_{xx} & \sigma_{xy} & 0 \\ \sigma_{xy} & \sigma_{yy} & 0 \\ 0 & 0 & \sigma_{zz} \end{pmatrix}. \quad (6)$$

For convenience, we define the following ratios:

$$\gamma = \frac{\sigma_{xx}}{\sigma_{yy}} \quad \& \quad \mu_0 = \frac{\sigma_{xy}}{-\sigma_{yy}} \quad (7)$$

which leads to:

$$\tan 2\Psi = -\tan 2\theta = -\frac{2\sigma_{xy}}{\sigma_{xx} - \sigma_{yy}} = \frac{2\mu_0}{\gamma - 1}. \quad (8)$$

Let σ_1 , σ_2 and σ_3 represent the maximum, intermediate and minimum compressive principal stresses. For a strike-slip fault, σ_1 and σ_3 should lie on the x - y plane and σ_2 should be parallel to the

z -axis, that is out of plane for a 2-D simulation (Fig. 1). Hence,

$$\begin{aligned} \sigma_1 &= \left(\frac{\sigma_{xx} + \sigma_{yy}}{2} \right) - R = \left(\frac{\sigma_{xx} + \sigma_{yy}}{2} \right) \left(1 + \sqrt{\left(\frac{\gamma - 1}{\gamma + 1} \right)^2 + \frac{4\mu_0^2}{(\gamma + 1)^2}} \right), \\ \sigma_3 &= \left(\frac{\sigma_{xx} + \sigma_{yy}}{2} \right) + R = \left(\frac{\sigma_{xx} + \sigma_{yy}}{2} \right) \left(1 - \sqrt{\left(\frac{\gamma - 1}{\gamma + 1} \right)^2 + \frac{4\mu_0^2}{(\gamma + 1)^2}} \right), \\ \sigma_2 &= \sigma_{zz} = \rho g z (1 - \lambda), \end{aligned} \quad (9)$$

where λ is the pore pressure coefficient and z is a parameter corresponding to the depth of the 2-D slice where plane strain simulations are conducted. Under the plane-strain approximation, we also have the following relationship:

$$\sigma_2 = \sigma_{zz} = \nu(\sigma_{xx} + \sigma_{yy}) = \nu\sigma_{yy}(\gamma + 1), \quad (10)$$

where ν is the Poisson's ratio.

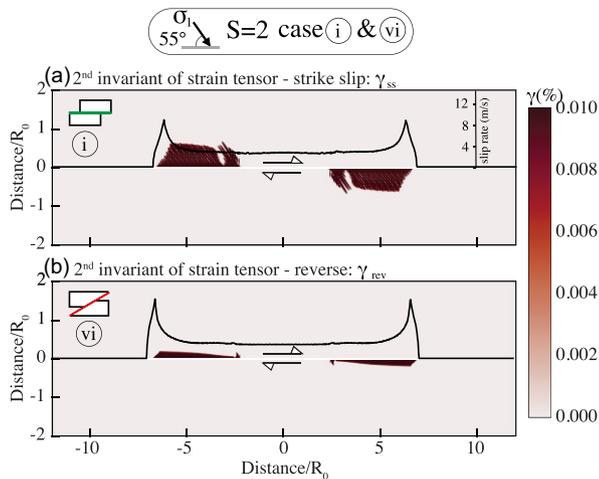


Figure 9. Plastic deformation for the models with $\Psi = 55^\circ$, $S = 2$, with initial strike-slip stress field for case (i) and reverse stress field for case (vi). Here we show the second invariant of the deviatoric strain tensor γ . Slip rate on the fault (black curve) is superimposed on the graphs.

We require that both σ_1 and σ_3 should be compressive, that is $\sigma_1 < \sigma_3 < 0$. This implies that

$$\sqrt{\left(\frac{\gamma-1}{\gamma+1}\right)^2 + \frac{4\mu_0^2}{(\gamma+1)^2}} < 1 \quad (11)$$

since $\sigma_{xx} + \sigma_{yy} < 0$. The principal stress state must also satisfy the inequality given in eq. (1).

$$\begin{aligned} \sigma_1 < \sigma_2 &\Rightarrow \left(\frac{\sigma_{xx} + \sigma_{yy}}{2}\right) \left(1 + \sqrt{\left(\frac{\gamma-1}{\gamma+1}\right)^2 + \frac{4\mu_0^2}{(\gamma+1)^2}}\right) < \nu(\sigma_{xx} + \sigma_{yy}) \\ &\Rightarrow \sqrt{\left(\frac{\gamma-1}{\gamma+1}\right)^2 + \frac{4\mu_0^2}{(\gamma+1)^2}} > 2\nu - 1. \end{aligned} \quad (12)$$

This is trivially satisfied as ν is always smaller than 0.5.

$$\begin{aligned} \sigma_3 > \sigma_2 &\Rightarrow \left(\frac{\sigma_{xx} + \sigma_{yy}}{2}\right) \left(1 - \sqrt{\left(\frac{\gamma-1}{\gamma+1}\right)^2 + \frac{4\mu_0^2}{(\gamma+1)^2}}\right) > \nu(\sigma_{xx} + \sigma_{yy}) \\ &\Rightarrow \sqrt{\left(\frac{\gamma-1}{\gamma+1}\right)^2 + \frac{4\mu_0^2}{(\gamma+1)^2}} > 1 - 2\nu. \end{aligned} \quad (13)$$

Thus, the stress field for a strike-slip fault must satisfy the following criterion:

$$1 - 2\nu < \sqrt{\left(\frac{\gamma-1}{\gamma+1}\right)^2 + \frac{4\mu_0^2}{(\gamma+1)^2}} < 1. \quad (14)$$

This inequality is plotted in Fig. 3 as a function of γ and μ_0 . The shaded areas define what we will from now on refer to as the ‘correct’ regime (strike-slip stress field). The white areas represent (1) the ‘forbidden’ regime (reverse faulting stress field) for which, this criterion is violated and (2) the stress field corresponding to a tensile regime. Superimposed on this graph are the initial parameters for several studies modelling strike-slip motion using plane-strain approximation. Three out of six studies used a far field stress field that favours reverse faulting, that is the out-of-plane stress is σ_3 and not σ_2 as it should be (Fig. 3a). If the modelling is only performed on a 2-D plane, like the vast majority of the published studies, the fault will still have a strike-slip motion even if this condition is not satisfied simply because the motion is restricted along one plane. However, the whole stress field would favour reverse faulting,

which will impact any model of off-fault deformation as we will demonstrate soon.

In Fig. 3(c), we plot the same criterion as a function of Ψ and the seismic ratio S , as defined by (Andrews 1976; Das & Aki 1977),

$$S = \frac{\mu_s(-\sigma_{yy}) - \sigma_{xy}}{\sigma_{xy} - \mu_d(-\sigma_{yy})} = \frac{\mu_s - \mu_0}{\mu_0 - \mu_d}, \quad (15)$$

where μ_s and μ_d correspond to the static and dynamic coefficient of friction, respectively, for a slip weakening law. S helps determine whether a rupture in 2-D, under homogeneous conditions, can reach supershear velocities ($S < 1.77$), or remains sub-Rayleigh ($S > 1.77$). For a fixed value of Ψ , fulfilling the criterion for strike-slip faulting strongly depends on μ_d . Interestingly, setting up a proper strike-slip stress field in agreement with Anderson theory, that is $\Psi \simeq 30^\circ$ for $\mu_s = 0.6$, requires supershear parameters, and this holds true for a large range of realistic μ_d values.

3 METHODS

3.1 Numerical model setup

In this study, we explore the boundary between the forbidden regime and the correct regime by modelling rupture on a 1-D right-lateral fault in a 2-D medium under plane-strain approximation (Fig. 2). We particularly focus on the influence of the initial stress conditions on off-fault stresses and inelastic deformation. We use the 2-D spectral element code SEM2DPACK (Ampuero 2012). Rupture propagation along the fault plane is governed by a slip-weakening friction law (e.g. Palmer & Rice 1973). Slip occurs when the on-fault shear stress reaches the shear strength $\tau_f = \mu^*(-\sigma_{yy})$. The friction coefficient μ^* depends on the cumulated slip (δ) and drops from a static μ_s to a dynamic μ_d value over a characteristic distance (D_c). The rupture is artificially nucleated within a patch where the initial shear stress σ_{xy} is set just above the fault strength (Fig. 2). Further on, since the difference in nucleation duration t_{nuc} between two models has no physical meaning, we will shift all the results in time so $t_{\text{nuc}} = t_0$. Following Kame *et al.* (2003), the minimum nucleation size L_c determined by the energy balance for a slip weakening law is:

$$L_c = \frac{16}{3\pi} \frac{\mu G}{(\sigma_{xy} - \tau_r)^2}, \quad (16)$$

where G is the fracture energy, defined as:

$$G = \frac{1}{2} D_c (\tau_p - \tau_r), \quad (17)$$

and where $\tau_p = \mu_s(-\sigma_{\text{eff}})$ and $\tau_r = \mu_d(-\sigma_{\text{eff}})$ are, respectively, the peak and residual stresses.

We further use the process zone size R_0 for a quasi-stationary crack, to normalize the length scales in our results. It corresponds to the length over which the friction drops, with ongoing slip, from the peak strength to the residual strength. Following Day *et al.* (2005) it is given by,

$$R_0 = \frac{9\pi}{32(1-\nu)} \frac{D_c \mu}{(\mu_s - \mu_d)(-\sigma_{yy})}. \quad (18)$$

In our models, R_0 is approximately equal to 1.0 km for $\Psi = 55^\circ$ and 1.6 km for $\Psi = 15^\circ$ (see Section 3.2 for an explanation on the parameters). We use a resolution of 30 m to ensure that the problem is correctly resolved. In order to scale the problem, all the modelled faults have a length of $30R_0$.

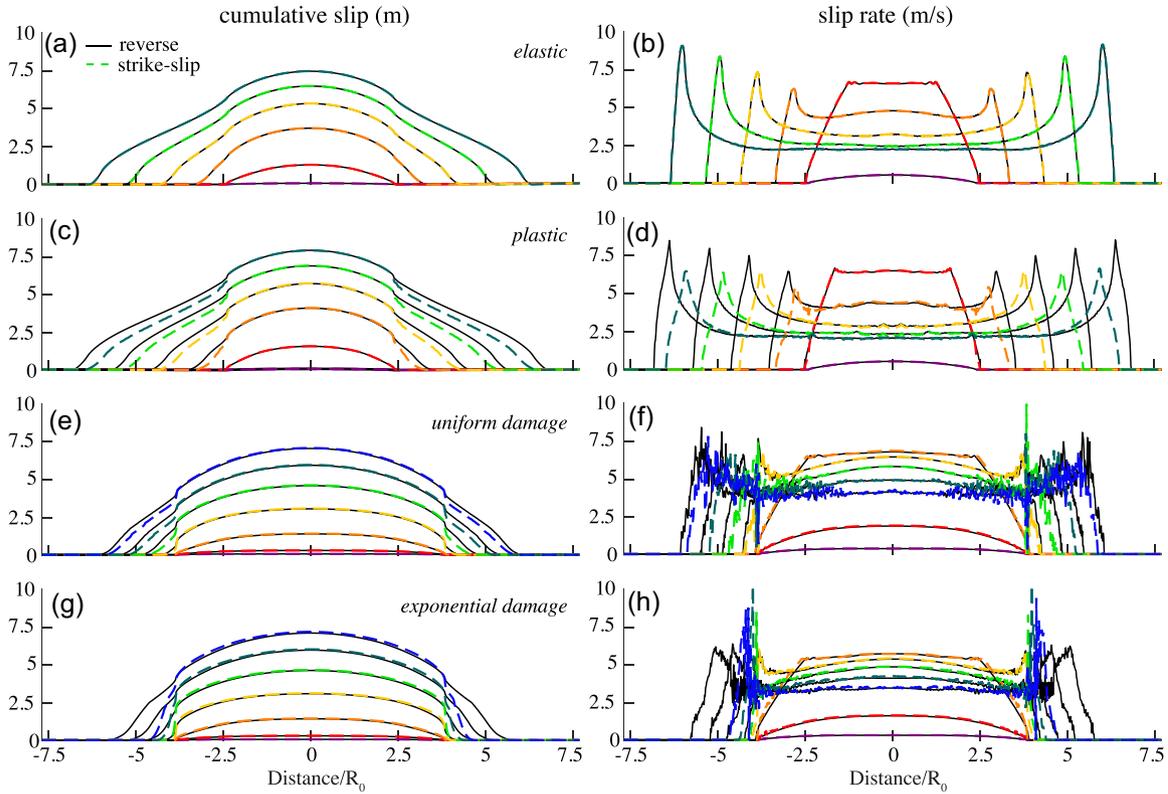


Figure 10. Cumulative slip (a-c-e-g) and slip rate (b-d-f-h) plotted every 0.4 seconds for $\Psi = 55^\circ$ and $S = 2$, with initial reverse stress fields (dark solid lines, case *vi*) and initial strike-slip stress field (dotted colour lines, case *i*).

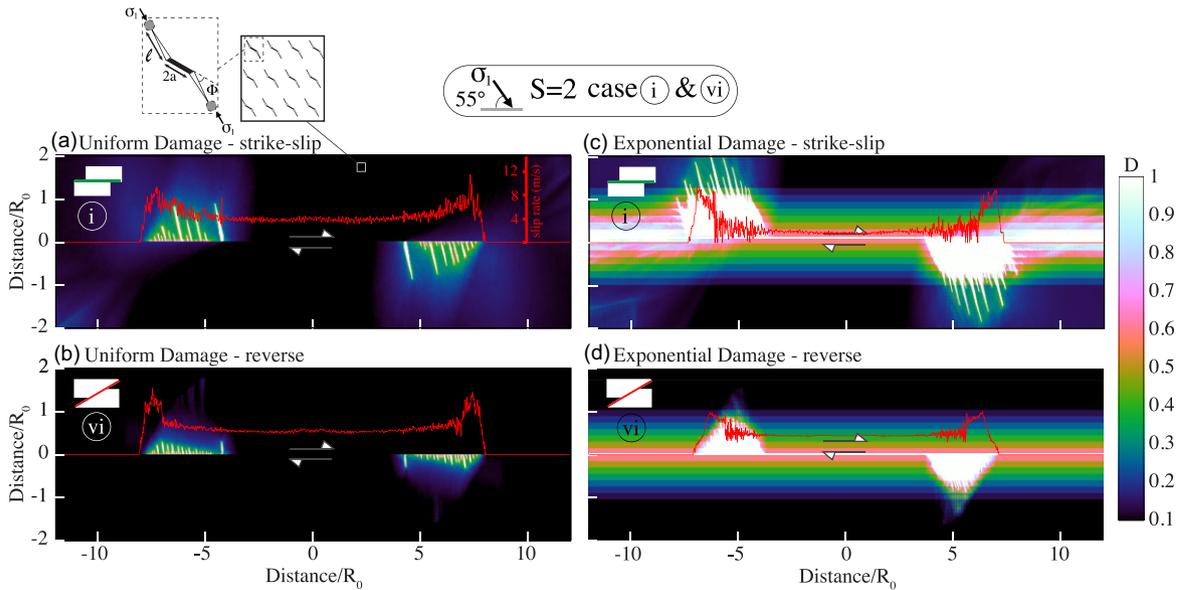


Figure 11. Simulation of a dynamic rupture with off-fault damage within a homogenous medium (a,b) and within a medium a pre-existing damage zone, with an exponential decay of damage density away from the fault (c,d). We explore the two end-member cases for $S = 2$ and $\Psi = 55^\circ$: case (*i*) corresponds to an initial strike-slip stress field (a,c) and cases (*vi*) to an initial reverse stress field (b,d). The colours represent the density of microcracks in the medium. Slip rate on the fault (red curve) is superimposed on the graphs. The inset shows a schematic representation of the initial crack distribution and of the crack geometry (after Thomas *et al.* 2017b).

3.2 Initial parameters

In order to compare our study to the available literature, we run two sets of models with $\Psi = 55^\circ$ and $\Psi = 15^\circ$, respectively (Fig. 3a). To explore the effect of initial stresses on off-fault and on-fault

deformation, for each set of models, we run three simulations with the correct strike-slip set-up (cases *i*, *ii*, *iii* for $\Psi = 55^\circ$ and cases *vii*, *viii*, *ix* for $\Psi = 15^\circ$) and three within the so-called forbidden regime (cases *iv*, *v*, *vi* for $\Psi = 55^\circ$ and cases *x*, *xi*, *xii* for $\Psi = 15^\circ$).

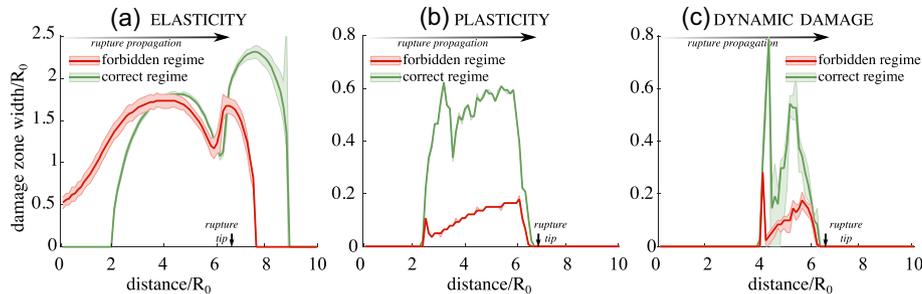


Figure 12. Synthesis of the effect of initial stress on damage zone width for cases (i) and (vi) with $\Psi = 55^\circ$ and $S = 2$, when the rupture has reached seven times the static process zone size. Results from the different models are displayed in red for case (vi), with reverse pre-stress conditions, and in green for case (i), with strike-slip pre-stress conditions. (a) Expected damage zone width using purely elastic models. We use the Coulomb stress change with a threshold of $\Delta CFF = 0.3\Delta\tau$. (b) Damage zone width computed with the model following Andrews (2005), using a threshold value of plastic deformation of 0.01 per cent. (c) Damage zone width computed with the micromechanical model, using a threshold value of $D = 0.3$ (medium with a uniform initial damage density $D_0 = 0.1$). The shaded areas correspond to ± 10 per cent changes in the threshold value.

To achieve that goal, we adopt the following strategy. Some parameters are kept constant between all the simulations: the elastic properties, the depth, the characteristic slip in the friction law and a hydrostatic pore pressure condition (Table 1). We have two Ψ that define the orientation of maximum compressive stress σ_1 with respect to the fault. We set the S ratio (eq. 15), favouring a value that would lead to subshear rupture ($S = 2$) for $\Psi = 55^\circ$. However, for $\Psi = 15^\circ$, this lead to a small dynamic stress drop which prevents the rupture from propagating (dying cracks). Hence we choose to set the ratio to $S = 1$ for $\Psi = 15^\circ$ (the rupture can evolve to a supershear earthquake). We further fix the stress drop (~ 10 MPa):

$$\Delta\tau = (\mu_0 - \mu_d)(-\sigma_{yy}) \quad (19)$$

and, in agreement with laboratory values (Jaeger 1979), we make μ_s varies between 0.52 and 0.68 to get the six models defined above.

The others parameters can be determined using the following set of equations. In hydrostatic pore pressure condition, the vertical, out-of-plane stress σ_{zz} is given by

$$\sigma_{zz} = \rho gz(1 - \lambda) \quad (20)$$

We then need to compute the ratio γ :

$$\begin{aligned} \text{eq.7\&10} &\Rightarrow \gamma = \frac{\sigma_{zz}}{v\sigma_{yy}} - 1 \\ \text{eq.19\&15} &\Rightarrow \gamma = \frac{\sigma_{zz}(\mu_0 - \mu_s)}{vS\Delta\tau} - 1 \\ \text{eq.8} &\Rightarrow \gamma = \frac{\sigma_{zz}(0.5(\gamma - 1)\tan 2\Psi - \mu_s)}{vS\Delta\tau} - 1 \\ &\Rightarrow \gamma = \frac{(\alpha + \beta) + \mu_s\sigma_{zz}}{(\beta - \alpha)} \\ &\text{with } \alpha = vS\Delta\tau \text{ \& } \beta = 0.5\sigma_{zz}\tan 2\Psi \end{aligned} \quad (21)$$

Knowing γ , we can derive:

$$\mu_0 = \frac{(\gamma - 1)}{2} \tan 2\Psi \quad (22)$$

$$\mu_d = \mu_0 - \frac{\mu_s - \mu_0}{S} \quad (23)$$

$$\sigma_{yy} = \frac{\sigma_{zz}}{v(\gamma - 1)} \quad (24)$$

$$\sigma_{xx} = \gamma\sigma_{yy} \quad (25)$$

$$\sigma_{xy} = \mu_0(-\sigma_{yy}) \quad (26)$$

The values of initial parameters for the end-member models are summarized in Table 2 while the corresponding Mohr–Coulomb

diagrams illustrating their initial stress fields can be found in the supplementary materials (see Fig. S1). For a fixed angle Ψ , note that because we set S constant between the six models, a constant stress drop is equivalent to a constant fracture energy (eq. 17). Therefore, the characteristic length scales for the friction law, R_0 and L_c , are also constant.

4 RESULTS

4.1 Role of pre-stresses in determining the yield criterion

In this section, we run simulations with an elastic medium and we compute different yield criterion commonly used in the literature to determine the inelastic deformation. It is important to note that in those cases, the medium has a pure elastic behaviour and the off-fault plastic deformation is calculated *a posteriori*.

4.1.1 Role of the out-of-plane stress in computing the plastic yield criterion

We first examine the importance of accounting for the out-of-plane stress in determining the off-fault plastic deformation. Here we compute the Drucker–Prager criterion:

$$F_{DP} = \sqrt{J_2} + \mu_s p, \quad (27)$$

where $p \equiv I_1/3 = (\sigma_{xx} + \sigma_{yy} + \sigma_{zz})/3$ is the hydrostatic stress derived from the first invariant I_1 of the stress tensor and $J_2 = s_{ij}s_{ij}/2$ corresponds to the second invariant of the deviatoric stress tensor (with $s_{ij} = \sigma_{ij} - p\delta_{ij}$). However, in many of the published 2-D studies (Templeton & Rice 2008; Dunham *et al.* 2011b, among others) the out-of-plane stress, σ_{zz} , is assumed to be the mean of the in-plane stresses, that is $\sigma_{zz} = (\sigma_{xx} + \sigma_{yy})/2$. This particular choice of σ_{zz} makes the Mohr–Coulomb and Drucker–Prager yield surfaces coincide in 2-D. Therefore, the invariants can be computed using the in-plane stress tensor components, that is $p^{ip} = (\sigma_{xx} + \sigma_{yy})/2$ which consequently changes the second invariant of the deviatoric stress tensor as well (further referred as J_2^{ip}). We note the Drucker–Prager criterion, solely using the in-plane stresses as follow:

$$F_{DP}^{ip} = \sqrt{J_2^{ip}} + \mu_s p^{ip}. \quad (28)$$

To illustrate the contribution of the out-of-plane stress in computing the invariants we use case (i) with $\Psi = 55^\circ$ and $S = 2$ (see Table 2). We compare the hydrostatic stress and the square root of

the second invariant of the deviatoric stress tensor obtained with and without the out-of-plane stress (Figs 4a and b) when the fault has ruptured about five times the process zone R_0 . We can observe that in both cases, the hydrostatic stress is higher in the compressional quadrants. However, if the out of plane stress is ignored, the hydrostatic stress is overestimated (stresses are positive in tension) by up to 20 MPa (or two times the dynamic stress drop) as displayed in Fig. 4(c). On the other hand, the second invariant of the deviatoric stress tensor is underestimated if we ignore σ_{zz} (Figs 4d–f). The difference between J_2 and J_2^{ip} is up to ~ 1.5 times the dynamic stress drop. This results in a significant difference in the estimation of the Drucker–Prager criterion (Figs 5a–c). In both cases, F_{DP} or F_{DP}^{ip} are positive in the tensional quadrants. However, taking into account the out-of-plane stress not only increases the area where the Drucker–Prager criterion is positive, that is where the plastic deformation is expected, but the overall magnitude of the plastic strain is also higher. Hence this leads to an underestimation of the plastic deformation. As expected, Fig. S2 reveals consistent findings when computing the Drucker–Prager criterion, solely using the dynamic change of stresses due to the rupture propagation.

Changing the initial state of stress, using case (vii), that is for $\Psi = 15^\circ$ and $S = 1$ (see Table 2) we obtained even more drastic differences. When the out-of-plane stress is ignored (Fig. 5e) F_{DP}^{ip} is pretty much negative everywhere, which may lead to the interpretation that no plastic deformation is happening. Whereas, when accounting for σ_{zz} , even if the magnitude is about four times smaller than for case (i), we record positive F_{DP} , notably within the compressional quadrants.

4.1.2 Role of the reverse versus strike-slip stress field on the plastic yield criterion

Now that we have emphasized the importance of including the full stress tensor in computing the Drucker–Prager criterion, we look at the influence of setting up a proper strike-slip initial stress field, ensuring that the pre-stress conditions satisfy $\sigma_1 < \sigma_2 = \sigma_{zz} < \sigma_3$.

In Fig. 6, for $\Psi = 55^\circ$, $S = 2$ and for $\Psi = 15^\circ$, $S = 1$ we compare the Drucker–Prager criterion of the end-member models, that is case (i), with case (vi) and case (vii) with case (xii). The pre-stress conditions for cases (i) and (vii) correspond to a strike-slip fault, whereas cases (vi) and (xii) have the initial stress field of a reverse fault. The yield criterion is computed when the fault has ruptured about five times the process zone R_0 . Note that, due to the difference in S ratio for the simulations with $\Psi = 55^\circ$ and $\Psi = 15^\circ$, it is not meaningful to compare the models with different angles, as the dynamic rupture is very different, and by extent, the off-fault deformation.

For initial reverse stress conditions, when $\Psi = 55^\circ$ (Fig. 6b, case vi), positive F_{DP} is observed in the tensile quadrants. When $\Psi = 15^\circ$ (Fig. 6e, case xii), positive F_{DP} is observed in the compressive quadrants, with a magnitude lower than for case (vi). When the initial stress field is properly set-up (Figs 6a and d), the areas recording a positive F_{DP} are much larger and the absolute value is also higher, as illustrated by Figs 6(c) and (f). Since the models are run within a purely elastic medium and since the rupture dynamics are very similar, the differences are essentially linked to the background stresses.

Fig. 7 shows the continuum of all models for $\Psi = 55^\circ$. Unlike Fig. 6 for which we plot F_{DP} at one particular time step, here we plot the maximum value of F_{DP} induced in the off-fault medium by the full rupture. Crossing the boundary between initial strike-slip

stress field and reverse pre-stress conditions does not change the results qualitatively. However, the area with positive F_{DP} is larger in the strike-slip case, and we observe higher absolute values, with up to twice the dynamic stress drop.

4.1.3 Role of the initial stress field for off-fault rupture modes

We have shown that the initial stress conditions influence the values of F_{DP} criterion induced by the rupture. Yet the Drucker–Prager criterion uses the invariants of the stress tensor, which does not predict the preferential orientation for failure, a useful information in seismic risk assessment for example. Another criterion often used in the literature is therefore the Coulomb stress change due to the rupture on the main fault (King *et al.* 1994; Stein *et al.* 1997; Thomas *et al.* 2017a; Canitano *et al.* 2021).

$$\Delta CFF = (\Delta\tau_y + \mu_s \Delta\sigma_{\text{eff}}), \quad (29)$$

where $\Delta\sigma_t$ and $\Delta\sigma_n$ are the change of shear and normal stress, respectively, on the optimally oriented plane, induced by the seismic rupture for a given direction θ with respect to the principal stress σ_1 (in relation to the Mohr–Coulomb circle displayed in Fig. 1). In Fig. 8, we explore the same cases that in Fig. 6, that is the two end-members for $\Psi = 55^\circ$, $S = 2$ (cases i and vi) and for $\Psi = 15^\circ$, $S = 1$ (cases vii and xii). The background colour corresponds to local values of ΔCFF . The yield criterion is computed when the fault has ruptured about five times the process zone R_0 . As observed previously with the Drucker–Prager criterion, areas with positive Coulomb stress change are larger and with higher ΔCFF values for $\Psi = 55^\circ$ (Figs 8a and b). Likewise, the areas likely to induce off-fault deformation are larger when the initial stress-field is properly set up, and this for the two tested values of Ψ (Figs 8a and c).

On top of these snapshots, we compute the local preferential orientations for failure and the corresponding type of faulting induced. An off-fault strike-slip failure means that locally, σ_1 and σ_3 are in-plane. If σ_3 is out-of-plane, the local preferred type of failure will be that of a reverse fault. Fig. 8(a) shows that if the initial strike-slip stress field is correctly set up, the rupture induces only strike-slip off-fault failures. However, if the initial stress field actually corresponds to that of a reverse faulting, the outcomes are different (Fig. 8b). The main rupture (strike-slip by default since we are in 2-D) only influences the stress field close to the rupture tips. Far from the fault and within the nucleation area, the propagating rupture does not control the local stress field, but its initial value does. Hence we observe off-fault reverse failures for case (vi).

4.2 Dynamic simulation with inelastic rheology

We have shown, using linear elasticity, that the pre-stress conditions can significantly affect the assessment of the different yield criteria used to estimate the off-fault deformation. In this section, we now investigate the role of the initial stress field on the dynamically triggered off-fault deformation, and its feedback on the seismic rupture, using different modelling strategies.

4.2.1 Role of the initial stress field in plastic deformation

We first use the off-fault plasticity model implemented in SEM2DPack following Andrews (2005). The inelastic response of the medium is characterized by a Coulomb criterion using in-plane stresses with $\mu = 0.75$ and $c = 30$ MPa so that the initial stress state

of the medium is below the yield criterion. The maximum shear stress over all orientations is :

$$\tau = \sqrt{\sigma_{xy}^2 + [(\sigma_{xx} - \sigma_{yy})/2]^2}. \quad (30)$$

The Coulomb criterion is :

$$\tau \leq c \cos \mu - 0.5(\sigma_{xx} + \sigma_{yy}) \cos \mu. \quad (31)$$

At each time step of the calculation, stress components are first incremented elastically. Then, if the Coulomb criterion is violated, stress components are recomputed so that a part of the deformation is accommodated inelastically (see Andrews 2005; Duan & Day 2008, for details about the method).

We compute the plastic deformation induced by the rupture for the two end-members cases (*i*) and (*vi*) previously studied ($\Psi = 55^\circ$ and $S = 2$). In Fig. (9), the results are illustrated by the second invariant of the deviatoric strain tensor γ_i :

$$\gamma_i = \sqrt{2e_{ij}e_{ij}}, \quad (32)$$

where $e_{ij} = \epsilon_{ij} - \frac{1}{3}\epsilon_{kk}\delta_{ij}$. The pattern of off-fault deformation is significantly different between the two simulations. The extent of the plastic deformation is much larger for an initial strike-slip stress state (Fig. 9a), as expected based on the results of Section 4.1. This induces differences in the rupture dynamics with a decrease of rupture speed and slip rate (Fig. 10d) and a lower cumulative slip (Fig. 10c), compared to case (*vi*) that has an initial stress field that corresponds to reverse faulting. Therefore, when off-fault inelastic deformation is taken into account, even modest differences in the initial stress condition affect significantly both the evolution of the off-fault medium and the slip dynamics.

We note that in a non-cohesive medium ($c = 0$), those differences are even more emphasized (Figs S3a and b). For initial strike-slip stress conditions, the rupture decelerates rapidly, its propagation is prevented by the intense inelastic deformation of the medium. The off-fault deformation is localized and optimally oriented with respect to the far field stress orientation. It is equivalent to the creation of a new optimally oriented fault. Hence, it is interesting to note that modelling off-fault deformation for a 2-D strike-slip fault, with the appropriate initial stress field, requires a certain cohesion in order to fully rupture the prescribed fault plane. We also note that for a non-cohesive medium, when $\Psi = 15^\circ$ and $S = 1$ (Figs S3c and d), little deformation occurs for the strike-slip case (*vii*) and none for the reverse case (*xii*). When $\Psi = 30^\circ$ (optimally oriented fault, Figs S3e and f), deformation only occurs on the main fault. Therefore, a flat fault has to be mis-oriented to produce significant off-fault deformation.

4.2.2 Role of the initial stress field on dynamic damage

The last set of simulations use a micromechanics-based model to determine the dynamically triggered off-fault damage and its feedback on the rupture dynamic. Inelastic deformation can occur in the model by either crack opening or crack propagation from initial flaws. Using an energy-based approach, at each time step, the corresponding change in elastic moduli, and hence the constitutive law, is determined (see Thomas *et al.* 2017b, for constitutive equations). The current inelastic state of the medium is characterized by the scalar D , the fraction of volume occupied by microcracks:

$$D = \frac{4\pi}{3} N_v (a \cos \Phi_c + l)^3, \quad (33)$$

where a is the initial microcrack radius, l , the wings cracks length as they are growing parallel to σ_1 ($l = 0$ at $t = 0$) (see Fig. 11) and N_v ,

the volume density of cracks. Initial flaws are all aligned at the same optimal angle $\Phi_c = \frac{1}{2} \tan^{-1}(1/\mu_c)$ to σ_1 , with μ_c being the friction coefficient for the microcracks. In their model, Bhat *et al.* (2012) derive the crack growth law by comparing the stress intensity factor at the microcrack tips to the experimentally-determined initiation and propagation toughness.

D varies between 0 and 1, the maximum value corresponding to the coalescence stage that leads to the macroscopic fracture of the solid. See Bhat *et al.* (2012) and Thomas *et al.* (2017b) for further details on the method.

Fig. 11 shows the damage density induced by the rupture for the end-member cases (*i*) and (*vi*) with $\Psi = 55^\circ$ and $S = 2$. The distribution of pre-existing flaws is homogeneous ($D = 0.1$ at $t = 0$) in Figs 11(a) and (b). For the models in Figs 11(c) and (d) we assume an exponential decay of initial damage with fault normal distance, as described in several field studies (e.g. Vermilye & Scholz 1998; Wilson *et al.* 2003; Mitchell & Faulkner 2009). The initial damage density varies from $D = 0.5$ to $D = 0.1$ over a distance equivalent to the process zone $R_0 \sim 1$ km. The characteristic scale was selected to match the width of the damage zone that may have been created by past earthquakes, in this case defined in Fig. 11(a). In all scenarios, to prevent off-fault damage at the beginning of the simulation due to far field loading, we set the friction on the microcracks, μ_c to 0.75. Hence the observed damage is dynamically triggered by the seismic rupture.

Similar to the models discussed earlier, for these particular stress states, damage essentially occurs in the tensional quadrants. The damage zone is also significantly wider for initial strike-slip stress conditions (Figs 11a and c). Previous studies, in comparison to simulations with a pure elastic medium, have underlined the effect of damage on slip rate and rupture velocity (slow down) and to a lesser extent the cumulative slip (Thomas *et al.* 2017b; Thomas & Bhat 2018). Here, when comparing in Fig. 10 the models with initial reverse stress fields (dark solid lines) and strike-slip stress field (dotted colour lines), we observe differences in cumulative slip (e and g), slip rate and rupture velocity (f and h), both for models with an uniform medium or an initial damage zone. That is because, unlike for the reverse case, when the initial stress field is correctly set up, damage occurs ahead of the rupture tip, thus changing the P - and S -wave speeds in the medium, which ultimately slows down the rupture velocity. This effect is even more pronounced when the earthquake ruptures a fault with a pre-existing damage zone (Figs 10g and h).

5 DISCUSSION AND CONCLUSIONS

As discussed above, in the last two decades, the wealth of observations have underlined the importance of off-fault deformation and complex structure in fault zone behaviour. Numerical models have been developed to incorporate these key ingredients (Andrews 2005; Shi & Ben-Zion 2006; Templeton & Rice 2008; Dunham *et al.* 2011b; Thomas *et al.* 2017b; Okubo *et al.* 2019, among others). However, due to numerical limitations, they have been developed mostly in 2-D. As a consequence, when setting up the initial stress field, only the in-plane stresses are defined and the out of plane stress is often ignored, or assumed to be the mean of the in-plane stresses.

In several studies, the effect of initial stress state have been investigated in terms of the orientation of σ_1 with respect to the fault direction. They illustrated the significant influence of Ψ on the pattern of off-fault damage (Poliakov *et al.* 2002; Rice *et al.* 2005;

Ngo *et al.* 2012; Templeton & Rice 2008). However, the impact of the relative importance of the principal stresses, at constant angle Ψ , has not been discussed. In this study, we run 2-D plane-strain simulation of a strike slip faulting to illustrate the key role of the 3-D faulting regime on off-fault deformation.

Using an elastic medium, we first show that, even if the initial stress field is rightfully set up, ignoring the out-of-plane stress (here σ_{zz}) leads to an underestimation of the inelastic deformation, both in extent and magnitude. If the initial stress field is on top wrongly defined (reverse faulting), ignoring σ_{zz} will on contrary lead to an overestimation of the inelastic deformation. Then, we demonstrate that a modest change in the pre-stress conditions, from strike-slip to reverse stress field, strongly influences the magnitude of any plastic criterion and the extent of off-fault deformation. Using a Coulomb criterion, we also compute the local preferential orientations for failure and the corresponding type of faulting induced. We have shown that a simulation within the so-called ‘forbidden regime’ will predict some reverse faulting in the off-fault medium.

Then, because of the various feedbacks that exist between the dynamic rupture and the bulk, we show that the discrepancy is even more pronounced when inelastic deformation can occur in the medium as the rupture propagates (Fig. 12). Both in plastic and dynamic damage models, the resulting pattern of inelastic deformation is significantly different. We show that a modest change in pre-stress conditions from initial reverse to initial strike-slip stress field would underestimate the damage zone width by a factor of 3–6 (Figs 12b and c). We would like to underline that, while previous numerical studies have modelled inelastic deformation under different stresses regimes (Shi & Ben-Zion 2006; Dunham *et al.* 2011b; Templeton & Rice 2008; Okubo *et al.* 2019), as illustrated in Fig. 3(a), they did not investigate the effect of pre-stress by keeping S ratio, stress drop, and angle Ψ simultaneously constant. By keeping these rupture parameters constant among our models, we demonstrate the importance of initial stress field only. We also show that the effect of far field stresses can have a significant impact on the rupture dynamics (Fig. 10). In a passive, elastic medium, pre-stress indeed does not affect fault slip. In inelastic medium, such as displayed in Figs 9 and 11 the evolving medium through energy loss and/or trapped-waves influences back both the slip rate and rupture velocity on the fault. We observe that this effect is more predominant with higher amount of off-fault deformation, hence when the initial stress field is rightfully set up and/or if a pre-existing damage zone is modelled. Therefore in 2-D plane-strain simulations, the initial 3-D stress field is important to model an accurate evolution of the off-fault medium.

To conclude, pre-stresses can significantly affect both off-fault damage and on-fault rupture dynamics even if other key parameters are kept constant: cohesive zone, nucleation size, seismic ratio, stress drop. Although none of the presented numerical models are meant to reproduce field observations exactly the shear increase in observations opens up the potential for statistical comparisons between models and observations. This makes it even more urgent to set-up the correct initial ‘3-D’ stress field even in ‘2-D’ numerical simulations.

SUPPORTING INFORMATION

Supplementary data are available at [GJI](https://doi.org/10.1093/gji/ggab001) online.

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DATA AVAILABILITY

This study uses numerical data only. All of the models and data sets are produced by the authors. The numerical model used to perform the simulations (Ampuero 2012) is available at the following link <https://github.com/jpampuero/sem2dpack>. The module developed by Thomas *et al.* (2017b) to simulate dynamic damage will be shared on request to the corresponding author. We provide a Matlab code to check a given initial stress field and plot Fig. 3 https://github.com/louisejeandetriebes/PreStress_2DplaneStrain.git

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